

**OLD SHERGOTTITES AND YOUNG IMPACT AGES**

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Abundant isotopic data obtained on basaltic shergottites for different long-lived radiogenic systems (<sup>87</sup>Rb-<sup>87</sup>Sr, <sup>147</sup>Sm-<sup>143</sup>Nd, <sup>176</sup>Lu-<sup>176</sup>Hf, U-Pb) give internal isochron ages that cluster around 180 and 330-475 Ma [1]. The general view that these ages represent the time of shergottite crystallization was questioned by Blichert-Toft et al. [2] and more recently by Bouvier et al. [3]. We argue that (i) regardless of the uncertainties, cratering chronology [4] and stratigraphy [5] assign to most of the Martian surface an age >2 Ga; (ii) evidence of old Pb-Pb ages in basaltic shergottites, for which the labile component (dominated by apatite) has been removed by acid leaching, is overwhelming; (iii) the preservation of <sup>142</sup>Nd and <sup>182</sup>W anomalies [6, 7] in shergottites is at odds with young ages; and (iv) other indications of old ages for Martian meteorites exist, such as 2.5-4.0 Ga old fission tracks in phosphate grains [8] and <sup>39</sup>Ar-<sup>40</sup>Ar ages in excess of emplacement ages [9, 10].

New Pb isotopic compositions of leached whole-rock fragments and maskelynite separates of Shergotty and Los Angeles fall on the ~4.1 Ga Zagami isochron of [3], which we interpret as dating the crystallization of the basaltic shergottite suite. In contrast, Sm-Nd and Lu-Hf internal isochrons give, as expected, young ages, namely respectively, 170 ± 41 Ma (MSDW = 6.3) and 187 ± 89 Ma (MSDW = 5.1) for Shergotty and 178.8 ± 4.6 Ma (MSDW = 0.95) and 158 ± 14 Ma (MSDW = 0.97) for Los Angeles. Both the ~180 and ~450 Ma cluster ages are observed by Ar-Ar chronometry in Martian meteorites, as well as in ordinary chondrites [11], and recently also by U-Pb chronology of shocked [12] baddeleyite in shergottites [13]. We surmise that these isotopic systems were reset by strong impacts on the Martian surface, which in turn raised global rock temperatures to 200-400°C [14], thereby creating the conditions for permafrost to melt and induce aqueous circulation and alteration near the surface. Such a scenario is supported by the atmospheric-like δD values measured in shergottite apatites [15, 16]. Consequently, the young ages of shergottites date the last re-equilibration of phosphate (the main carrier of U, Sr, and REE) with Martian groundwater before the last dry-out of these water bodies.

**References:** [1] Nyquist, L.E. et al. 2001. *Space Sc. Rev.* 96: 105-164. [2] Blichert-Toft, J. et al. 1999. *Earth Planet. Sc. Lett.*, 173: 25-39. [3] Bouvier, A. et al. 2005. *Earth Planet. Sc. Lett.*, 240: 221-233. [4] Hartmann, W.K. and Neukum, G. 2001. *Space Sc. Rev.*, 96: 165-194. [5] Tanaka, K.L. et al. 1992. *Mars*, 345-382. [6] Kleine, T. et al. 2004. *Geochim. Cosmo. Acta*, 68: 2935-2946. [7] Foley, C.N. et al. 2005. *Geochim. Cosmo. Acta*, 69: 4557-4571. [8] Rajan, R.S. et al. 1986. *Geochim. Cosmo. Acta*, 50: 1039-1042. [9] Bogard, D.D. and Garrison D.H. 1999. *Meteorit. Planet. Sc.*, 34: 451-473. [10] Walton, E.L. et al. 2007. *Geochim. Cosmo. Acta*, 71: 497-520. [11] Bogard, D.D. 1995. *Meteoritics*, 30: 244-268. [12] El Goresy, A. 1965. *J. Geophys. Res.* 70: 3454-3465. [13] Herd, C.D.K. et al. 2007. 38<sup>th</sup> LPSC, A1664. [14] Stöffler, D. et al. 1986. *Geochim. Cosmo. Acta*, 50: 889-903. [15] Leshin, L.A. 2000. *Geophys. Res. Lett.*, 27: 2017-2020. [16] Greenwood, J.P. et al. 2007. 38<sup>th</sup> LPSC, A1338.